

# Lecture 10

## part 2

# Slice method

### 1. Particle method (Lect. 8-9):

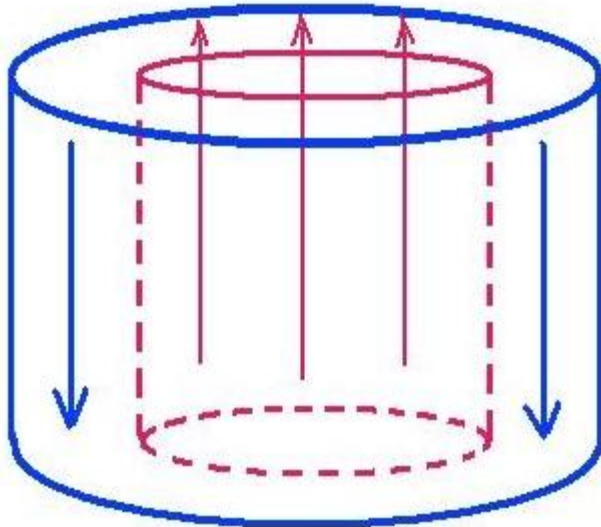
assumption of particle motion within motionless environment.

### 2. Basic approaches of slice method: 1938-1939 (Bjerknes J., Petersen S., Shishkin N. )

# Let's pick out a sufficiently large layer of air

$S'$  cross section(inner)

$$w' > 0$$



$z_2$  is upper layer

$z_1$  is lower layer

$S''$  cross section(outer)

$$w'' < 0$$

# Slice method assumptions:

- all variations of the quantities within the picked out layer come out adiabatically;
- no advection (horizontal mass motion) is observed;
- the mass of air above any layer remains unchanged.

The latter means that fluxes of air mass through section  $S'$  and  $S''$  are equal:

$$\rho' S' w' = -\rho'' S'' w'' \quad (10.8)$$

$\rho$  is density

$$\rho' \approx \rho''$$

$$S' w' = -S'' w'' \quad (10.9)$$

# To derive the stability criterion:

$\Delta T$  is temperature difference,

$T'$  is updraft flux temp.,

$T''$  is the downdraft one at  $z$  level

$$\Delta T = T' - T''$$

$\Delta T > 0$  unstable state (positive acceleration)

$\Delta T < 0$  stable state (negative acceleration)

Suppose, particles come from lower level to upper one:

$$T' = T_1 - \gamma' (z - z_1) \quad (10.10)$$

$$T'' = T_2 + \gamma'' (z_2 - z) \quad (10.11)$$

$\gamma'$  is adiabatic lapse rate in updraft flux

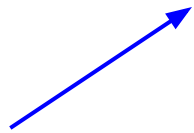
$\gamma''$  is adiabatic lapse rate in downdraft flux

$$z - z_1 = w' \Delta t \quad (10.12)$$

$$z_2 - z = -w'' \Delta t \quad (10.13)$$

From (10.10)-(10.13) we obtain:

$$\Delta T = T_1 - T_2 - (\gamma' w' - \gamma'' w'') \Delta t \quad (10.14)$$



$$T_1 - T_2 = \gamma (z_2 - z_1) = \gamma (z_2 - z) + \gamma (z - z_1) = \gamma \Delta t (w' - w'')$$

$$\Delta T = \left[ \gamma' (w' - w'') - (\gamma' w' - \gamma'' w'') \right] \Delta t \quad (10.15)$$

# Let's use (10.9)

$$S' w' = -S'' w'' \quad (10.9)$$



$$\Delta T = \left[ (\gamma' (w' - w'') - (\gamma' w' - \gamma'' w'')) \right] \Delta t \quad (10.15)$$

$$\Delta T = \left[ (\gamma - \gamma') + \frac{S'}{S''} (\gamma - \gamma'') \right] w' \Delta t \quad (10.16)$$



# Let's consider 3 cases:

1.  $\gamma' = \gamma'' = \gamma_a$

- Updraft and downdraft air isn't saturated

$$\Delta T = (\gamma - \gamma_a) \left(1 + \frac{S'}{S''}\right) w' \Delta t \quad (10.17)$$

- Criteria of stability of both methods (parcel and slice ones) give the same result.

## 2<sup>nd</sup> case:

$$2. \gamma' = \gamma'' = \gamma'_a$$

- Updraft and downdraft air fluxes are saturated

$$\Delta T = (\gamma - \gamma'_a) \left(1 + \frac{S'}{S''}\right) w' \Delta t$$

- Criteria of stability of both methods (parcel and slice ones) give the same result.

## 3<sup>d</sup> case:

$$3. \gamma' = \gamma'_a, \gamma'' = \gamma_a$$

- Updraft air is saturated and downdraft flux is not.

$$\Delta T = \left[ (\gamma - \gamma'_a) + \frac{S'}{S''} (\gamma - \gamma_a) \right] w' \Delta t \quad (10.18)$$

- Cu clouds development

- From (10.18)  $\Delta T > 0$  at  $\gamma > \gamma_a$  and

$$\Delta T < 0 \quad \text{at } \gamma < \gamma'_a$$

- Stability criteria of both methods are coincide if the stratification is absolutely unstable and absolutely stable. However, often

$$\gamma'_a < \gamma < \gamma_a$$

$$\Delta T \leftarrow \text{ on } \gamma \text{ but also } \frac{S'}{S''}$$

# Let's introduce a *critical lapse rate*

$$\gamma_{cr} \text{ at } \Delta T = 0$$

$$\left[ (\gamma - \gamma_a') + \frac{S'}{S''} (\gamma - \gamma_a) \right] w' \Delta t = 0$$

$$w' \neq 0, \quad \Delta t \neq 0$$

$$\left[ (\gamma - \gamma_a') + \frac{S'}{S''} (\gamma - \gamma_a) \right] = 0$$

$$\gamma_{cr} = \frac{\gamma_a' + \gamma_a (S'/S'')}{1 + S'/S''} \quad (10.19)$$

$\gamma_{cr} \rightarrow \gamma'_a$  at small values of  $(S'/S'')$

$\gamma_{cr} \rightarrow \gamma_a$  at large values of  $(S'/S'')$

$$\gamma > \gamma_{cr} \quad (\Delta T > 0)$$

$$\gamma < \gamma_{cr} \quad (\Delta T < 0)$$

$$\frac{S'}{S''} < \frac{\gamma - \gamma'_a}{\gamma_a - \gamma} \quad \text{at } \Delta T > 0 \quad (10.20)$$

$$\frac{S'}{S''} < \frac{\gamma - \gamma'_a}{\gamma_a - \gamma} \quad \text{at } \Delta T < 0 \quad (10.21)$$

- at conditionally **unstable** stratification the atmosphere is unstable for **small** size particles and it's **stable** for **LARGE** size particles =>>the atmosphere is said to be **selectively unstable**